

1 Article

2 Machine Learning–Based Reconstruction of Missing Stream- 3 flow Records in Data-Scarce River Basins

4 Muhammad Faarid Shah^{1*}, Muhammad Farrukh¹, Fahad Shamraiz¹, Touqeer Ali Rind¹, Farjad Aziz¹, Asad Wahab¹
5 and Qazi Khurshid Ahmed¹

6 1 Department of Civil Engineering (DCvE), Ghulam Ishaq Khan (GIK) Institute of Engineering Sciences and
7 Technology, Topi, Khyber Pakhtunkhwa; muhammadfarrukhkhajjak@gmail.com, [qureshifaha-
9 had9999@gmail.com](mailto:qureshifaha-
8 had9999@gmail.com), touqeer.ali@giki.edu.pk, gcv2568@giki.edu.pk, gcv2520@giki.edu.pk,
10 gcv2445@giki.edu.pk

* Correspondence: gcv2560@giki.edu.pk

11 Abstract

12 Stable streamflow records are essential in water-resources planning, flood prediction, and
13 effecting climatic change; when plentiful, however, in high-altitude and trans-boundary
14 river basins because of rugged terrain, severe climatic conditions, and insufficient moni-
15 toring facilities. The paper fills in missing streamflow records in the Chitral River Basin, a
16 snow- and glacier-fed basin of the Upper Indus Basin, by reconstructing missing dis-
17 charge at the downstream Arandu station with paradoxical observations of the Chitral
18 station. The statistical and machine-learn models that were used to analyze daily dis-
19 charge data between 1981-2024 are Linear Regression, Artificial Neural Network (ANN),
20 and Extreme Gradient Boosting (XG Boost). The coefficient of determination (R^2), root
21 mean square error, mean squared error, and mean absolute error were used to assess how
22 well a model works. Although the predictive power of all models was quite robust with
23 R^2 exceeding 0.95, ANN showed the most balanced and stable performance at both train-
24 ing and testing stages with the lowest prediction errors and best generalization. On these
25 findings, the ANN model was used to recreate missing discharge at the Arandu station to
26 generate a long-run series of streamflow between 1981 and 2024. The restored hydrograph
27 maintains the seasonal snowmelt-based flow regime and fits well with observed records
28 to actually serve as a virtual discharge sensor. The suggested framework is a feasible ap-
29 proach to hydrological reconstruction in mountainous basins with limited data, as well as
30 the enhanced water-resources management in the context of variable climatic conditions.

31 **Keywords:** Streamflow reconstruction; Machine learning; Chitral River Basin.

33 1. Introduction

34 Access to sound hydrology data, especially streamflow data are one of the persistent
35 issues in most river basins across the globe. The data gaps usually occur because of the
36 extreme climatic conditions, rugged terrain, inadequate monitoring facilities and opera-
37 tional or political limitations, particularly where there are mountains and cross-border
38 areas. The missing streamflow records pose a serious constraint to this water resources
39 planning, flood forecasting and climate change impacts assessment. To overcome these
40 limitations, machine learning (ML) methods have been actively embraced as data-driven

41 approaches towards predicting and reconstructing missing hydrological information, being
42 more flexible and accurate than traditional statistical or empirical procedures.

43 Recent studies show that ML models have a high potential to be used in imputing
44 streamflow data in a wide variety of hydroclimatic contexts. Modern neural-network-
45 based methods, like the Probabilistic Fusion Imputer with Neural Networks (PROFINN),
46 have been incredibly successful in terms of reconstruction accuracy, averaging RMSE of
47 0.91 and NashSutcliffe Efficiency (NSE) of 0.93 in the Pamba River basin successfully val-
48 idating various flow regimes and data-gap situations [1]. The ensemble learning tech-
49 niques have also demonstrated solid performances. RF with clustering methods has re-
50 sulted in a mean NSE of more than 0.85 in Indian catchments [2]. On the same note, the
51 predictive capability of Gradient Boosting and Bagging Regressors has been reported to
52 be very high in the estimation of streamflow and the reported accuracy rates of the two
53 methods lie in the range of 0.9737 to 0.9968 [3]. With limited data, simpler ML models like
54 Naive Bayes and k-Nearest Neighbors (KNN) have performed better than traditional
55 methods, and the former has been significantly useful when training data are limited [4].

56 Although such improvements are made, the performance of ML-based streamflow
57 prediction is made very dependent on the availability of data, climatic variability, and
58 hydrological processes that are unique to the basin. Streamflow gaps can span several
59 months in areas where there is high hydroclimatic variability, which makes models more
60 challenging to train and predict. As a result, ML models should be properly adjusted to
61 local conditions, to guarantee the assurance of successful performance [2]. Recent research
62 also shows that combining ML techniques and physical hydro knowledge enhances the
63 reliability of predictions in ungauged and under-gauged basins, especially where the ter-
64 rain is complex [5].

65 The river basins with high altitudes are important in the freshwater cycle of the
66 planet since snow and ice stored in winter are melted during the spring and summer sea-
67 sons. A classic example is the Hindu Kush-Karakoram-Himalayan (HKH) area that is cen-
68 tral and plays a major water tower to South and Central Asia and is a provider of river
69 systems that supply almost 800 million people [6]. The temperature changes are very sen-
70 sitive to hydrological processes in this region, and directly determine the dynamics of the
71 snow melting and glaciers. Even minor shifts in temperature linked to climate change can
72 have a profound effect of the magnitude of streamflow and the distribution of how it
73 changes throughout the year.

74 The Upper Indus Basin (UIB) has streamflow regimes that are significantly controlled
75 by seasonal snowmelt and glacier melt. According to historical studies, there are trends
76 towards surface and subsurface hydrology, at a rate of about 0.03-0.05 times per minute
77 per second, which mainly is caused by the intensified melting of glaciers, changes in pre-
78 cipitation levels, and variations in seasonal climate cycles [7]. The trends are extremely
79 dangerous to future water supply and overall, to the possibility of extreme hydrological
80 processes [8].

81 The UIB and hydrological monitoring of tributaries are still inadequate as topogra-
82 phy is rugged, climatic conditions extreme, and several different river systems are trans-
83 boundary. This has led to the fact that streamflow records are usually incomplete or not
84 available, especially in the upstream and border regions. This has been linked to the well-
85 known ungauged or under gauged catchment problem that poses significant problems to
86 hydrological modeling, flood forecasting and water resources management [8]. Conven-
87 tional gap-filling methods, including interpolation or transposition of flows within neigh-
88 boring stations are frequently not suitable to snow- and glacier-fed rivers, due to their
89 poor ability to model complicated flow behavior. This has led to the emergence of region-
90 alization methods that make use of hydrological similarity and upstream-downstream

linkages in the transfer of information to poorly measured sites and those better measured, which improves the reliability of data in the mountainous basin [9].

2. Materials and Methods

2.1 Study Area and Hydrological setting

Chitral River Basin is a large mountainous branch of the Kunar-Kabul River system and an important part of the Kabul River Basin (KRB). Although this basin is strategically hydrologically significant to the downstream water supply, it has received relatively little coverage in scientific literature specifically in long-term flow reconstructions and climate sensitivity.

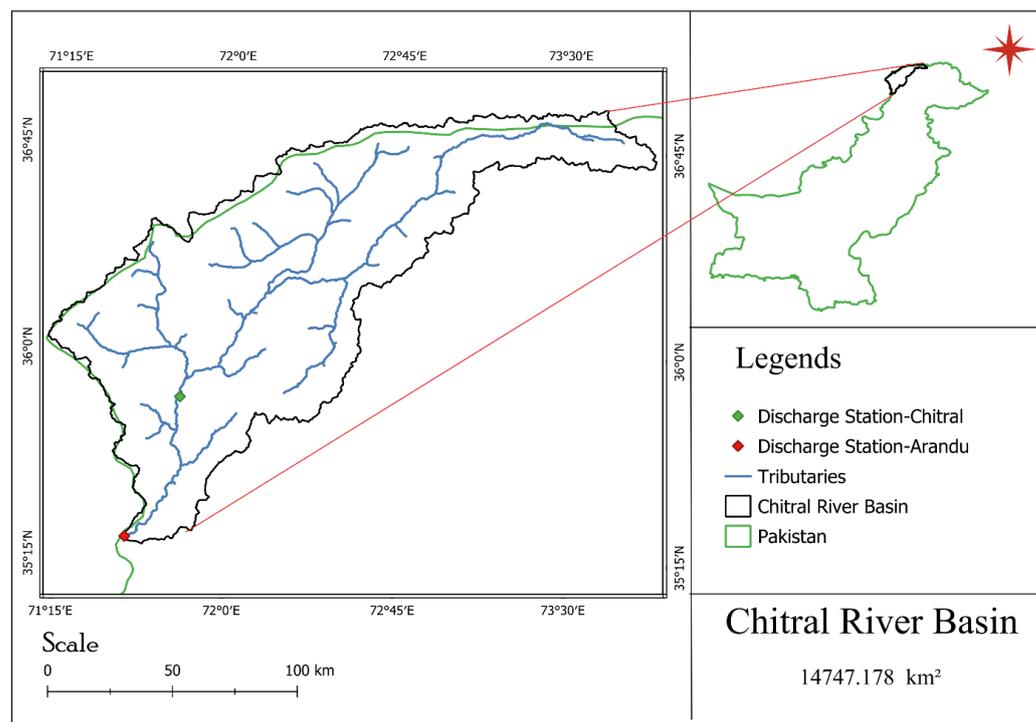


Figure 1: Chitral River Basin

The Chitral River is the one that originates on the mountain range of Hindu Kush and its catchment area is about 14,747km² (Figure 1). The basin is largely snow and glacier-driven and the streamflow is mostly controlled by seasonal snowmelt and, to a lesser degree, spring and summer precipitation. The outcome of these cryosphere controls is very strong non-linear hydrological behavior and intra-annual variability.

Monitoring in the basin is very minimal. The downstream Arandu gauging station located at the border area of Pakistan and Afghanistan (35°19'0" N, 71°34'0" E), installed in 2008 by WAPDA, is characterized by sporadic records that are short in duration thereby limiting sound interpretation of the long-term flow variability and water availability in the future. Conversely, the upstream Chitral station (35°51'48.0"N, 71°47'15.0"E) installed in 1963 by WAPDA possesses a rather long and continuous discharge history tracing back to the early sixties. This difference forms a significant problem to hydrological evaluation and infrastructural planning in the whole basin.

2.2 Machine Learning Applications

Recent developments in machine learning (ML) and hydro-informatics offer potentially useful tools to overcome such data scarcity. ANNs especially have proven to be highly capable of modeling streamflow in data-constrained mountainous catchments be-

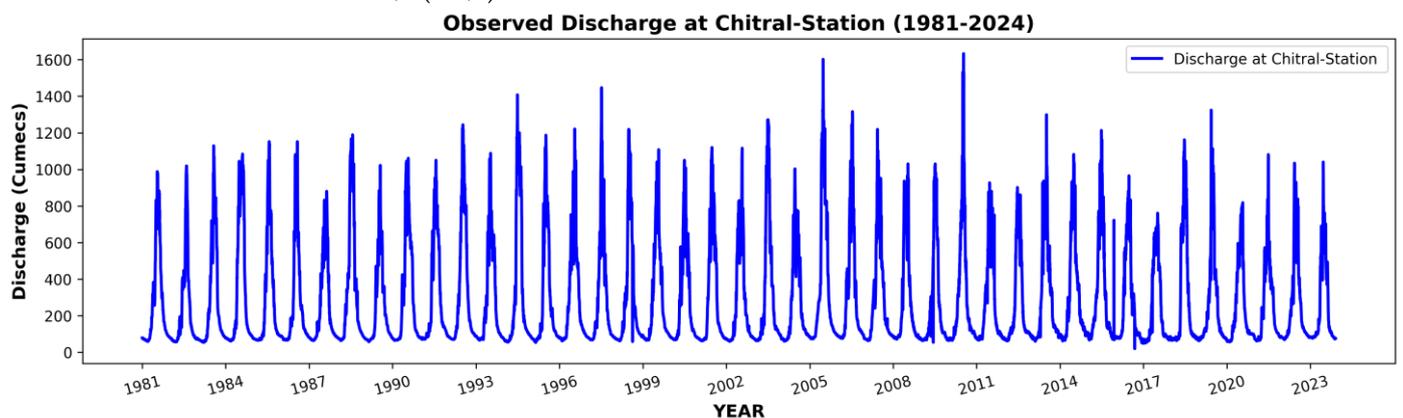
119 cause they can model even complex and non-linear relationships without necessarily being
 120 parameterized. Past research has indicated that under those circumstances, the ML-
 121 based methods often work better than the conventional autoregressive and conceptual
 122 models [10].

123 Although there is an increase in the use of ML to reconstruct medium-term stream-
 124 flow even at smaller high-altitude sub-basins within the Hindu Kush-Himalaya (HKH)
 125 and UIB, like the Chitral-Arandu system, the long-term reconstruction of streamflow is
 126 limited. Majority of the current studies are performed on the large rivers, as small tribu-
 127 taries are underrepresented [7]. Such incomplete long-term discharge information deteri-
 128 orates water resources planning, design of hydraulic infrastructure and evaluation of cli-
 129 matic-change effects in these sensitive areas [11].

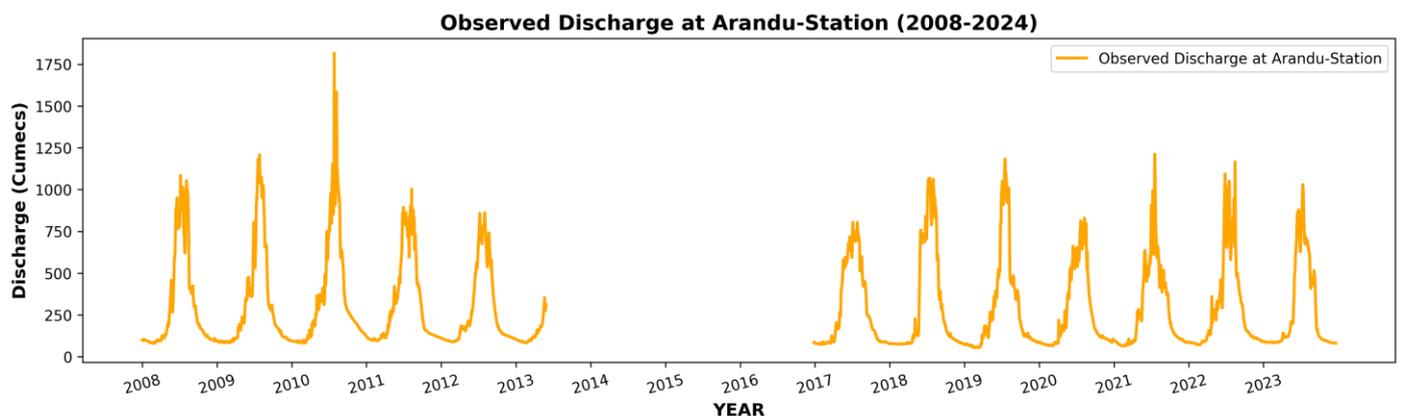
130 To fill this gap, the current paper uses almost 43 years of hydrological data (1981-
 131 2024) to recreate streamflow in the Arandu station using both machine-learned and sta-
 132 tistical methods. Linear Regression (LR) and models such as ANN, Support Vector Re-
 133 gression (SVR), random forest (RF), Decision Trees (DT) and Extreme Gradient Boosting
 134 (XG Boost) are used.

136 2.3 Data Collection and Study Period

137 Two hydrometric stations on the basin were used to acquire data on daily discharge,
 138 namely Chitral and Arandu located at (35°19'0" N, 71°34'0" E) and (35°51'48.0"N,
 139 71°47'15.0"E). The Chitral station has a continuous discharge record since 1981 to 2024,
 140 and the Arandu station has continuous discharge record since 2008 to 2024, although some
 141 years are missing, 2013-2016. At the two stations, discharge is reported through cubic me-
 142 ters/s (m³/s).



143 **Figure 2:** Observed discharge at Chitral Station (1980-2024)



145 **Figure 3:** Observed Discharge at Arandu Station (2008-2024)

147 The Chitral station was chosen as the main predictor since it had a long and con-
148 sistent history and its hydrological interest to the downstream Arandu station. The two
149 stations share similar seasons flow regimes, due to melting snow and glaciers, which im-
150 plies a high level of hydrological connectivity between the upstream and the downstream
151 areas as reported in earlier research [2]. This connectivity provides the reason why up-
152 stream discharge is an explanatory factor in the ways of reconstituting downstream flows.

153 2.4 ML Models

154 .In this study, three machine-learning models of varying complexity were used to
155 forecast the missing streamflow data at the downstream station. Linear Regression (LR),
156 Artificial Neural Networks (ANN) and Extreme Gradient Boosting (XGBoost) were cho-
157 sen to serve as a representation of the linear, non-linear and the ensemble-based methods
158 of learning, respectively. This is because the application of a variety of models enables
159 them to assess their comparative performance in data-limited and highly varying hydro-
160 logical conditions.

161 2.4.1 Linear Regression (LR)

162 LR was used in order to develop a direct linear correlation between upstream and
163 downstream discharge as a baseline model. Due to its simplicity and interpretability, LR
164 offers a reference framework in which the added value of machine-learning approaches
165 could be evaluated. Nevertheless, it is not applicable in cases when hydrological processes
166 are highly un-linear and this is typical of snow and glacier-fed river systems [12].

167 2.4.2 Artificial Neural Networks (ANN)

168 ANNs were used to learn non-linear input/output relationships that do not explic-
169 itly parameters like physical ones. The feed-forward type of network was adopted because
170 ANNs are highly applicable to the modeling of complex hydrological systems that are
171 affected by the cryosphere processes. They have been found to be highly predictive in
172 streamflow modeling and especially in low-data situations [12],[13].

173 2.4.3 Extreme Gradient Boosting (XG Boost)

174 XG Boost was used as an ensemble method of learning that constructs sequential de-
175 cision trees with the aim of reducing errors in prediction via gradient boosting. XG Boost
176 is computationally efficient and it is able to deal with non-linear relationships, heteroge-
177 neous inputs and missing values. The latter features are what make it applicable to model
178 mountain streamflow that is highly variable and prone to adverse effects of noise and
179 limited observations on the simpler models.

181 3.5 Model Evaluation

182 There were four statistical indicators that were used to evaluate the model perfor-
183 mance: the coefficient of determination (R^2), root mean square error (RMSE), mean
184 squared error (MSE) and mean absolute error (MAE). These measures are regularly used
185 in hydrological modeling to measure the correspondence between observed and modeled
186 streamflow as well as to gauge the predictive capacity within distinct flow circumstances.

187 With coefficient of determination (R^2), the proportion of variance in the observed dis-
188 charge that is explained by the model predictions is measured. It gives a measure of good-
189 ness of fit of a model and capability to reproduce temporal variability in streamflow but
190 fails to give an actual measure of the magnitude of prediction error [14].

191 Root mean square error (RMSE) is a square root of the mean of squared errors of the
192 observed and predicted values. RMSE is also keen on huge deviations, and thus it puts
193 more emphasis on the errors in peak-flow, hence it is especially applicable in examining
194 the performance of the model in high-flow [15].

The mean squared error (MSE) is the average of squared differences in the observed and simulated discharge. It gives a direct account of total prediction error, and is typical as an objective term when calibrating a model, but is highly susceptible to outliers as it squares errors [16].

Mean absolute error (MAE) is used to compute the average value of absolute differences between predicted and observed values. MAE is also less susceptible to extreme values than RMSE and MSE and provides a powerful estimate of typical error of predictions in highly varying hydrological time series [14].

3. Results

3.1. Model Performance in Training

Linear regression (LR), Artificial Neural Networks (ANN) and XG Boost were initially tested on the training set. Table 1 is a summary of training performance measures, whereas Figure 4 shows the correlation between the inflow (Chitral) and the predicted outflow (Arandu) of all three models.

Table 1: Training Data Evaluation Metrics

Model	R ²	RMSE	MSE	MAE
Linear Regression	0.95	63.6	4045.3	36.16
ANN	0.95	63.91	4084.56	34.56
XG Boost	0.96	55.06	3031.93	31.56

Train Data: Inflow vs Outflow Prediction

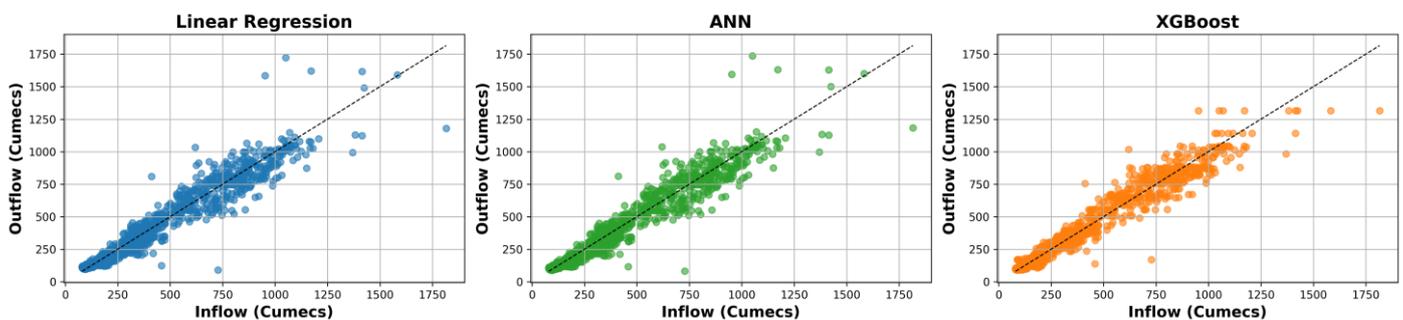


Figure 4: Training Data's Scattered Plots

All models revealed a high level of agreement between observed and simulated discharge during the training with coefficients of determination (R²) more than 0.95. Linear Regression was able to model downstream discharge variability with an R² of 0.95 and thus a large percentage of the variability could be attributed to a linear relationship between upstream inflow. Nevertheless, the values of RMSE (63.6 m³/s) and MAE (36.16 m³/s) are relatively high, indicating lower accuracy in a high-flow regime.

The ANN model had a similar explanatory power (R² = 0.95) but the error properties, especially MAE (34.56 m³/s), were better, which represents the average discharge conditions better. The ANN scenario plot shows a higher concentration around the 1:1 line than that of LR particularly at medium-to-high flows.

XG Boost performed better in the training phase as it gave a high R² (0.96) and a low RMSE (55.06 m³/s), MSE (3031.93) and MAE (31.56 m³/s). The training scatter plot shows that it has good alignment with observed discharge throughout the entire flow range implying that it has a good fitting capability.

3.2 Model Performance in Testing

The generalization of the model was checked on the independent testing dataset (Table 2, Figure 5). The three models were all very predictive, which proved the strength of the relationship between upstream and downstream discharges.

Table 2: Testing Data Evaluation Metrics

Model	R ²	RMSE	MSE	MAE
Linear Regression	0.97	47.24	2231.4	34.75
ANN	0.97	45.48	2068.74	31.24
XGBoost	0.96	55.13	3039.29	38.7

Test Data: Inflow vs Outflow Prediction

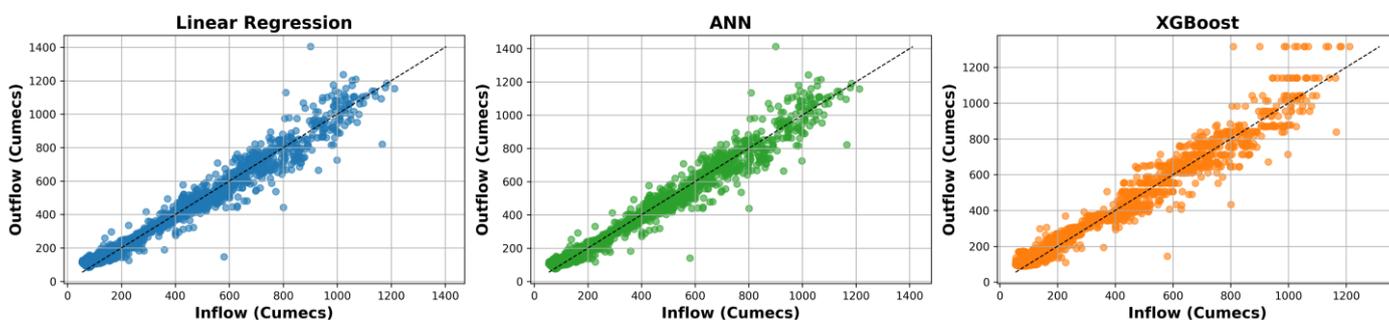


Figure 5: Testing Data's Scattered Plots

Linear Regression resulted in R² of 0.97, RMSE and MAE of 47.24 m³/s and 34.75 m³/s respectively. Although the high R² is an indicator of good overall agreement, this dispersion was more at higher flows implying that the 1:1 line could not capture non-linear variability effectively.

ANN model proved to be the most balanced in terms of performance in the course of testing as it was the most accurate among all models. The lowest RMSE (45.48 m³/s), MSE (2068.74) and MAE (31.24 m³/s), were observed. Testing scatter plot indicates less bias and narrows the clustering around the 1:1 line than LR and XGBoost especially under peak flow conditions.

To the contrary, XGBoost showed a deterioration in performance when tested. Even though the R² was still large (0.96), the measures of errors have worsened significantly (RMSE = 55.13 m³/s; MAE = 38.7 m³/s). The higher the scatter and deviation with increased discharges, the more evidence that there was some overfitting in the training and that it would no longer be able to generalize.

3.3 Choosing the most effective model.

Despite the better results of XG Boost in training (Table 1), its lower accuracy in testing implies that it cannot be used for independent prediction. The assumption that Linear Regression is stable is true but it indicated a much higher error (Table 2) and was not able to represent non-linear hydrological behavior.

The ANN model proved that the most consistent and reliable performance was observed both during the training and testing sessions, and the explanatory power is high,

and the total prediction errors are small. ANN was chosen as the best model in the reconstruction of missing streamflow data at the Arandu station due to its superior generalization capability as well as balanced error structure.

3.4 Missing Discharge at Arandu Station Reconstruction (1981-2024).

The trained ANN model was used to recreate missing discharge values in the Arandu station by referring to the long-term upstream station discharge history in the Chitral station. The full reconstructed discharge series of Arandu in 1981-2024 is shown in, which is a combination of measured values and ANN-based predictions.

Figure

6

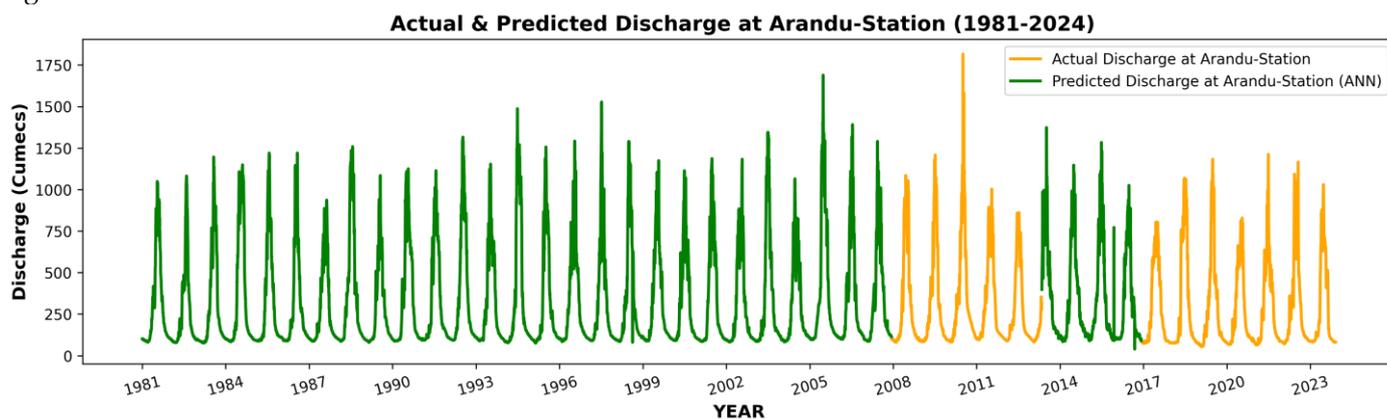


Figure 6: Actual and Predicted discharge at Arandu (1981-2024)

The reconstituted hydrograph maintains the seasonal snowmelt-dominated flow regime within the basin, and annual peaks in the spring and summer and low flows in the winter. The ANN-based estimates blend seamlessly with the observed records of discharges, meaning that there is consistency in terms of time and time-respecting plausibility. This re-created long-term data is, in effect, a virtual discharge sensor, which has been extended by over 20 years of useful Arandu data

4. Discussion

The findings of this work have validated that machine-learning models can be used as a useful tool in recreating missing records of streamflow in snow- and glacier-based mountain basins where traditional hydrological surveillance practices are constrained. The high coefficients of determination throughout all the models confirm that there is a high hydrological relationship between the upstream Chitral and downstream Arandu stations, which supports the validity of upstream-downstream information transfer in the basin. Nevertheless, the distinction between model behavior is observed when metrics of error and generalization are analyzed.

Though relatively simple and easy to interpret, Linear Regression demonstrated greater prediction inaccuracies and dispersion during medium-high flows, which demonstrates its inability to capture the non-linear nature of the processes related to cryosphere melting. XG Boost performed well in terms of fitting in the training phase and poor in accuracy in the testing phase, indicating that it is partially overfitting and not that robust when it is used on independent data. The sensitivity of complex ensemble models to variability of hydroclimatic conditions and small sample sizes in mountainous settings is emphasized by this behavior.

Conversely, the ANN model exhibited a high predictive ability both in training and testing with less magnitudes of error and closer fitting about the 1:1 line especially at the

296 peak flows. This balanced performance depicts a higher generalization and stability, and
297 ANN is therefore more applicable in long term streamflow reconstruction. The recon-
298 structed series at Arandu of reconstructed discharges retains the snow melting regime
299 which is seasonally controlled and has temporal consistency with observed discharges,
300 which confirms its physical plausibility. The usage of reconstructed series as a virtual hy-
301 drometric station would offer a viable way to expand streamflow records in under-
302 gauged parts of the Upper Indus Basin and other high-altitude catchments.

303 5. Conclusions

304 This paper shows that machine-learning methods can be used to reconstruct lost
305 streamflow data as in the Chitral River. Using the high hydrological interrelations be-
306 tween upstream and downstream stations, a reconstruction of the missing discharge of
307 the Arandu station was achieved over a 43-year span. Though all the analyzed models
308 demonstrated a high level of explanatory power, the Artificial Neural Networks contin-
309 ued to outperform Linear Regression and XG Boost in regard to making predictions and
310 their capacity to generalize. The reconstruction of the ANN based reconstruction provided
311 a continuous and physical realistic discharge series that maintains the snowmelt domi-
312 nated seasonality of a basin. The long dataset, in effect, serves as a virtual discharge sensor
313 and will play a major role in improving availability of long-term hydrologic data on the
314 basin. The proposed framework has the benefit of providing a transferable and firm solu-
315 tion to close the data divide in other under-gauged mountain systems and transboundary
316 rivers to aid in better hydrological evaluation, infrastructure design, and analysis of im-
317 pact of climate changes.

318 6. Patents

319 No Patents have resulted from the work reported in this manuscript.

320 Author Contributions:

321 The conceptualization, methodology development, data collection, formal analysis, model
322 testing, and data validation, initial drafting and final pape preparation were carried out by Muham-
323 mad Faarid Shah, Muhammad Farrukh and Fahad Shamraiz. Touqeer Ali Rind played a key role in
324 the initial drafting, methodology development, and critical analysis of the results. Farjad Aziz con-
325 tributed to data validation and model implementation. Asad Wahab and Qazi Khurshid Ahmed
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327 All authors have read and agreed to the published version of the manuscript.

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331 **Informed Consent Statement:** Not applicable. This study did not involve human participants.

332 **Data Availability Statement:** The discharge data used in this study were obtained from hydromet-
333 ric stations operated by the Water and Power Development Authority (WAPDA), Pakistan. Re-
334 strictions apply to the availability of these data, which were used under permission for the current
335 study and are not publicly available. Processed data supporting the findings of this study may be
336 made available from the corresponding author upon reasonable request and subject to data-sharing
337 approvals.

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Abbreviations

The following abbreviations are used in this manuscript:

ANN	Artificial Neural Network
LR	Linear Regression
XG Boost	Extreme Gradient Boosting
UIB	Upper Indus Basin
MSE	Mean Squared Error
RMSE	Root Mean Squared Error
MAE	Mean Absolute Error

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